



Review on Estimation and Forecast of Reference Evapotranspiration

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The evapotranspiration process plays an essential role in the hydrological cycle, exerting a great influence on various aspects of life, including agricultural practices, water resource management, environmental assessments, and drought indices. However, many regions around the world lack reference sites or costly physical networks for direct evapotranspiration measurements, making estimation a vital criterion in these water-scarce areas. Over the years, researchers have employed empirical and/or physics-based models to compute reference evapotranspiration, a fundamental parameter for crop water requirements in diverse climatic regions. Moreover, forecasting reference evapotranspiration using climate data provides on-time information for irrigation scheduling and water resource management. Consequently, an array of techniques, from conventional statistical approaches to advanced deep learning techniques, have been employed by researchers to predict reference evapotranspiration. In this comprehensive review, a wide range of estimation and forecasting models for reference evapotranspiration is presented, encompassing both traditional techniques and cutting-edge forecasting methods.

Keywords: Evapotranspiration, Estimation of reference evapotranspiration, Forecasting of reference evapotranspiration

Introduction:

The term ‘Evapotranspiration’ was first defined by Thornthwaite in 1944 both actually and potentially, which is the major and largest element of the world water cycle. Evapotranspiration is the water loss from soil, open water surface, and crop surfaces to the atmosphere [1], [2], [3], [4], [5]. About 70% of the total global precipitation is accounted for by evapotranspiration, which consists of sublimation from snow packs, evaporation from open water surfaces and bare soil, and transpiration from vegetation through leaf stomata [6], [7]. Transpiration contributes 50% to 70% of the total evapotranspiration, which shows the impact of vegetation on overall evapotranspiration rates [8].

The rate of evapotranspiration that occurs if there is an unlimited and continuous supply of soil moisture is called potential evapotranspiration, or the theoretical rate of evapotranspiration in actual weather and vegetation conditions when soil moisture is not limiting (surface is saturated) Thornthwaite, [9]. Potential evapotranspiration is the empirical function of air temperature, day length, relative humidity, and wind speed. The rate at which the water from the ground and vegetation is evaporated to the atmosphere under actual conditions is called actual evapotranspiration. In a dry climate, the potential evapotranspiration is greater than the actual evapotranspiration, but in a rainy climate, the actual evapotranspiration is essentially equal to potential evapotranspiration [10], [11]. Considerable

irrigation is required for crops that need to transpire water at a closer rate to the potential evapotranspiration rate in those areas where precipitation is less than the potential evapotranspiration [9].

The maximum rate of evapotranspiration for a specific crop when the availability of water to the crop is such that the maximum transpiration is satisfied is called crop evapotranspiration [10], [12]. Crop water requirements are directly related to crop evapotranspiration which is the product of reference evapotranspiration and a crop coefficient [13]. Reference evapotranspiration is the theoretical evapotranspiration for a reference crop (short well-watered green grasses i.e., alfalfa) [14], [15]. Crop evapotranspiration can be measured directly for a specific crop, but field measurement of crop evapotranspiration is laborious, requires more time, and is costly, due to which crop evapotranspiration can be calculated by using evapotranspiration from a reference surface and crop coefficient [16], [17]. A two-step approach is usually used to obtain the crop evapotranspiration [18];

- i. Reference evapotranspiration is estimated which indicates the atmospheric demand of a reference crop at a substantial surface of growing green grass of the same height and having no shortage of water.
- ii. Reference evapotranspiration is then converted into crop evapotranspiration by adjusting with a crop coefficient which depends on the type of crop, growth stage, density, and soil moisture [10], [19].

Therefore, precise quantification of reference evapotranspiration is necessary for crop water requirements, hydrological studies, environmental assessment, water resource management, and drought indices [20], [21], [22]. Atmospheric parameters (air temperature, wind speed, radiation, and difference in vapor pressure), types of vegetation, water availability on the surface, and soil moisture control the rate of reference evapotranspiration [23].

Objectives:

The primary objective of this study is to comprehensively assess and synthesize existing research on the estimation and forecasting of reference evapotranspiration, with a focus on evaluating the strengths, limitations, and applications of various models.

Novelty:

In this study a wide range of estimation and forecasting models for reference evapotranspiration is presented, encompassing both traditional techniques and cutting-edge forecasting methods.

Estimation of Reference Evapotranspiration:

Directly reference evapotranspiration is measured either by using lysimeters [24], [25], [26], [27], Eddy Covariance Techniques [28], [29], [30], [31], [32], [33], [34], and Bowen Ratio Energy Balance System, or from a reference crop such as a perennial grass [18], [35]. However, direct measurement of reference evapotranspiration is not always possible due to the requirement of highly accurate experimental design and high cost [36], [37]. Therefore, alternatively, reference evapotranspiration can be computed using mathematical approaches based on empirical relationships among the climatic variables [10], [38]. Different temperature-based [9], [18], radiation-based [18], [39], mass transfer-based [40], [41], [42], [43], and combination-based [10], [44], [45] approaches are used to estimate reference evapotranspiration.

The globally accepted and most widely used method for the estimation of reference evapotranspiration is the Penman-Monteith FAO56 (Food and Agriculture Organization Paper-56) method due to its better performance [46], [47]. However, a large number of required meteorological variables often restrict its applicability in regions of poorly observed and/or limited meteorological data [10], [48], [49], [50], [51], [52]. The missing variables can be calculated by analytical methods; however, they may result in inaccuracy. Similarly, the applicability of the Penman equation can be highly variable from region to region due to the

non-availability of all the required meteorological parameters. As a result, it is necessary to have a substitute empirical method(s), which can be used for more precise computation of reference evapotranspiration in the area of interest with different input parameters and under different climatic conditions. Based on a literature survey it can be safely stated that most of the models used for estimation of reference evapotranspiration are given in Table 1. The abbreviation of variables, coefficients, and units is given in Table 2.

Table 1: Models for Estimation of Reference Evapotranspiration

S.No	Model Type	Model	Equation	Reference
1	Combination Based	Penman	$ET_o = \frac{(\frac{\Delta}{\Delta+\gamma})(R_n - G) + 6.43(\frac{\gamma}{\Delta+\gamma})(1 + 0.536u_2)(e_s - e_a)}{\lambda}$	[42]
2		Penman-Monteith	$ET_o = \left[\frac{\Delta(R_n - G) + K_{time}\rho_a c_p \frac{(e_s - e_a)}{r_a}}{\Delta + \gamma(1 + \frac{r_s}{r_a})} \right] / \lambda$	[53]
3		Wright-Penman	$ET_o = \frac{(\frac{\Delta}{\Delta+\gamma})(R_n - G) + 6.43(\frac{\gamma}{\Delta+\gamma})(a_w + b_w u_2)(e_s - e_a)}{\lambda}$ $a_w = 0.3 + 0.58 \exp \left[-\left(\frac{J - 170}{45} \right)^2 \right]$ $b_w = 0.32 + 0.54 \exp \left[-\left(\frac{J - 228}{67} \right)^2 \right]$ J is the Julian day of the year	[54]
4		Penman-Monteith FAO56	$ET_o = \frac{0.408\Delta(R_n - G) + \gamma \frac{900}{(T_{mean} + 273)} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)}$	[10]
5		Valiantzas	$ET_o = 0.051(1 - \alpha)R_s \sqrt{ T_{mean} + 9.5 } - 2.4\left(\frac{R_s}{R_a}\right)^2 + 0.048(T_{mean} + 20) \left(1 - \frac{RH_{mean}}{100}\right) (0.5 + 0.536u_2) + 0.00012z$ Where $\alpha = 0.25$	[55]
6			$ET_o = 0.0393R_s \sqrt{ T_{mean} + 9.50 } - 0.19R_s^{0.6} \left(\varphi \frac{\pi}{180}\right)^{0.15} + 0.048(T_{mean} + 20) \left(1 - \frac{RH}{100}\right) u_2^{0.7}$	[56]
7			$ET_o =$	

			$0.0393R_s\sqrt{ T_{\text{mean}} + 9.50 } - 0.19R_s^{0.6}(\varphi \frac{\pi}{180})^{0.15}$ $+ 0.0061(T_{\text{mean}} + 20)(1.12T_{\text{mean}} - T_{\text{min}} - 2)^{0.7}$	
8			$ET_o =$ $0.0393R_s\sqrt{ T_{\text{mean}} + 9.50 } - 0.19R_s^{0.6}(\varphi \frac{\pi}{180})^{0.15}$ $+ 0.078(T_{\text{mean}} + 20)(1 - \frac{RH_{\text{mean}}}{100})$	
9	Pan Evaporation Based	Singh	$ET_o = ET_{\text{Pan}}K_{\text{pan}}$ $K_{\text{pan}} = \frac{0.85(\Delta + \gamma)}{[\Delta + \gamma(1 + 0.33u_2)]}$	[57], [58]
10	Temperature Based	Blaney and Criddle	$ET_o = p(0.46T_{\text{mean}} + 8.13)$	[59]
11			$ET_o = 25.4 \frac{(1.8T_{\text{mean}} + 32)}{180} p$	[60]
12		Linacre	$ET_o = \frac{(\frac{700(T_{\text{mean}} \pm 0.006z)}{100 - \varphi}) + 15(T_{\text{mean}} - T_{\text{dew}})}{80 - T_{\text{mean}}}$	[61]
13		Modified Blaney and Criddle	$ET_o = a + b[p(0.46T_{\text{mean}} + 8.13)]$ $a = 0.0043RH_{\text{min}} - \left(\frac{n}{N}\right) - 1.41$ $b = 0.82 - 0.0041RH_{\text{min}} + 1.07\left(\frac{n}{N}\right) + 0.066u_{2d}$ $- 0.006RH_{\text{min}}\left(\frac{n}{N}\right) - 0.0006RH_{\text{min}}u_{2d}$	[14]
14		Hargreaves and Samani	$ET_o = [0.0023R_a(T_{\text{mean}} + 17.8)(T_{\text{max}} - T_{\text{min}})^{0.5}]/\lambda$	[39]
15		Trajkovic	$ET_o = [0.0023R_a(T_{\text{mean}} + 17.8)(T_{\text{max}} - T_{\text{min}})^{0.424}]/\lambda$	[62]
16		Tabari & Talaei-1	$ET_o = [0.0031R_a(T_{\text{mean}} + 17.8)(T_{\text{max}} - T_{\text{min}})^{0.5}]/\lambda$	[3]
17		Tabari & Talaei-2	$ET_o = [0.0028R_a(T_{\text{mean}} + 17.8)(T_{\text{max}} - T_{\text{min}})^{0.5}]/\lambda$	[3]
18		Droogers & Allen-1	$ET_o = [0.003R_a(T_{\text{mean}} + 20)(T_{\text{max}} - T_{\text{min}})^{0.4}]/\lambda$	[63]

19		Droogers & Allen-2	$ET_o = [0.0025R_a(T_{mean} + 16.8)(T_{max} - T_{min})^{0.5}]/\lambda$	[48]
20		Dorji	$ET_o = [0.002R_a(T_{mean} + 33.9)(T_{max} - T_{min})^{0.296}]/\lambda$	[64]
21		Baier & Robertson	$ET_o = [0.109R_a + 0.157 T_{max} + 0.158 (T_{max} - T_{min}) - 5.39]$	[65]
22		Ahooghala ndari-1	$ET_o = \left[0.252 \left(\frac{R_a}{\lambda} \right) + 0.221 T_{mean} (1 - RH_{\frac{mean}{100}}) \right]$	[66]
23		Ahooghala ndari-2	$ET_o = \left[0.29 \left(\frac{R_a}{\lambda} \right) + 0.15 T_{max} (1 - RH_{\frac{mean}{100}}) \right]$	[66]
24		Schendel	$ET_o = 16 \frac{T_{mean}}{RH}$	[63]
25		Ramazan Method	$ET_o = [(0.817 + 0.00022z)(0.0023R_a)(T_{mean} + 17.8)(T_{max} - T_{min})^{0.5}]/\lambda$	[67]
26		Modified Hargreaves	$ET_o = [0.00193R_a(T_{mean} + 17.8)(T_{max} - T_{min})^{0.517}]/\lambda$	[68]
27			$ET_o = 0.1555R_a(T_{max} - T_{min})^{0.5}(0.09967 + 0.00428T_{mean})$	[69]
28			$ET_o = 10^{-4}(0.002z + 7)(T_{mean} + 36.6)(T_{max} - T_{min})^{0.5}R_a$	[70]
29		Heydrari-Heydari	$ET_o = 0.0023(T_{mean} + 9.519)(T_{max} - T_{min})^{0.611} \cdot \frac{R_a}{\lambda}$	[71]
30		Hadria	$ET_o = 0.002(T_{mean} + 26.3)(T_{max} - T_{min})^{0.397} \cdot \frac{R_a}{\lambda}$	[72]
31		Samani	$ET_o = 0.0135[0.00185(T_{max} - T_{min})^2 - 0.0433(T_{max} - T_{min}) + 0.4023](T_{mean} + 17.8)(T_{max} - T_{min})^{0.5} \cdot \frac{R_a}{\lambda}$	[73]
32	Radiation Based	FAO-24 Radiation	$ET_o = (1.066 - 0.0013RH_{mean} + 0.045u_2 - 0.0002RH_{mean}u_2 - 0.000031RH_{mean}^2 - 0.0011u_2^2) \cdot \left(\frac{\Delta}{\Delta + \gamma} \right) \left(\frac{R_s}{\lambda} \right) - 0.3$	[18]

33		Makkink	$ET_o = 0.61 \left(\frac{R_s}{\lambda} \right) \left(\frac{\Delta}{\Delta + \gamma} \right) - 0.12$	[74]
34			$ET_o = 0.7 \left(\frac{\Delta}{\Delta + \gamma} \right) \left(\frac{R_s}{\lambda} \right)$	[75]
35		Modified Makkink	$ET_o = c \left[\left(0.25 + 0.50 \frac{n}{N} \right) R_a \right] \left(\frac{\Delta}{\Delta + \gamma} \right)$ where c is the adjustment factor and a function of mean humidity and daytime wind conditions	[18]
36		Priestly & Tayler	$ET_o = 1.26 \left(\frac{R_n - G}{\lambda} \right) \left(\frac{\Delta}{\Delta + \gamma} \right)$	[76]
37		Hargreaves	$ET_o = 0.0135 (T_{mean} + 17.8) (R_s / \lambda)$	[77]
38		Abtew-1	$ET_o = 0.52 T_{max} R_s / \lambda$	[78]
39		Abtew-2	$ET_o = (T_{max} / 56) (R_s / \lambda)$	
40		Irmak-1	$ET_o = -0.611 + 0.149 R_s + 0.079 T_{mean}$	[79]
41		Irmak-2	$ET_o = 0.469 + 0.289 R_n + 0.023 T_{mean}$	
42		Tabari & Talaei-3	$ET_o = -0.642 + 0.174 R_s + 0.0353 T_{mean}$	[3]
43		Tabari & Talaei-4	$ET_o = -0.478 + 0.156 R_s - 0.0112 T_{max} + 0.0733 T_{min}$	
44		Oudin	$ET_o = \left[\left(\frac{R_s}{\lambda} \right) (T_{mean} + 5) / 100 \right], \text{ (for } T_{mean} + 5 > 0)$	[80]
45		Jensen & Haise	$ET_o = (0.025 T_{mean} + 0.08) (R_s / \lambda)$	[81]
46		Jensen	$ET_o = C_T (T_{mean} - C_x) K_t R_a (T_{max} - T_{min})^{0.5}$ $K_t = 0.00185 (T_{max} - T_{min})^2 - 0.0433 (T_{max} - T_{min}) + 0.4023$ $C_T = \frac{1}{45 - \left(\frac{z}{137} \right) + \left(\frac{365}{e_{s,max} - e_{s,min}} \right)}$ $C_x = 22.5 - 0.14 (e_{s,max} - e_{s,min}) - \frac{z}{500}$ $e_{s,max} = \exp \frac{19.08 T_{max} + 429.41}{T_{max} + 237.3}$	[11]

			$e_{s,max} = \exp \frac{19.08T_{min} + 429.41}{T_{min} + 237.3}$	
47		Turc	$ET_o = \frac{0.013T_{mean}}{T_{mean}+15} \left(\frac{23.8856R_s+50}{\lambda} \right) \text{ for } RH \leq 50\%$ $ET_o = \left(1 + \frac{50-RH_{mean}}{70} \right) \frac{0.013T_{mean}}{T_{mean}+15} \left(\frac{23.8856R_s+50}{\lambda} \right) \text{ for } RH \geq 50\%$	[82]
48		Modified Turc	$ET_o = 0.013C_u(23.88R_s + 50) \frac{T_{mean}}{T_{mean} + 15}$ $C_u = -0.0211u_2^2 + 0.1109u_2 + 0.9004$	[83]
49		Turc (1961) with reduced wind data	$ET_o = (23.8856R_s + 50) \frac{0.013T_{mean}}{T_{mean} + 15} \left[1 + W_{RH} \left(0.71 - 1.43 \frac{RH}{100} \right) \right]$ $W_{RH} = 1 \text{ for } RH < 50\%, \text{ and } W_{RH} = 0 \text{ for } RH \geq 50\%$	[84]
50		Valiantzas	$ET_o = 0.00668R_a \sqrt{ T_{mean} + 9.5 } \cdot (T_{max} - T_{min}) - 0.0696(T_{max} - T_{min}) - 0.024(T_{mean} + 20) \left(1 - \frac{RH_{mean}}{100} \right) - 0.00455R_a(T_{max} - T_{dew})^{0.5} + 0.0984(T_{mean} + 17)(1.03 + 0.00055(T_{max} - T_{min})^2 - \frac{RH_{mean}}{100})$ $T_{dew} = \frac{116.91 + 273.3 \ln(e_a)}{16.78 - \ln(e_a)}$	[56]
51		Modified Bair-Robertson	$ET_o = 0.0039 T_{max} + 0.184 (T_{max} - T_{min}) + 0.1136R_a + 2.811(e_s - e_a) - 4$	[85]
52		Caprio	$ET_o = 0.1092708T_{mean} + (0.0060706)R_s$	[86]
53		Ritchie equation	$ET_o = \alpha_1 [3.87 \times 10^{-3} R_s (0.6T_{max} + 0.4T_{min} + 29)]$ <p>If</p> $5 < T_{max} \leq 35^\circ C, \text{ then } \alpha_1 = 1.1$ $T_{max} > 35^\circ C, \text{ then } \alpha_1 = 1.1 + 0.05(T_{max} - 35)$	[87]

			$T_{\max} < 5^{\circ}\text{C}$, then $\alpha_1 = 0.01 \exp [0.18(T_{\max} + 20)]$	
54		Copais Equation	$ET_o = 0.057 + 0.277C_2 + 0.643C_1 + 0.0124C_1$ $C_1 = 0.6416 - 0.00784RH_{\text{mean}} + 0.372R_s$ $- 0.00264R_s \cdot RH_{\text{mean}}$ $C_2 = -0.0033 + 0.00812T_{\text{mean}} + 0.101R_s$ $+ 0.00584R_s \cdot T_{\text{mean}}$	[88]
55		Castaneda-Rao	$ET_o = 0.7 \left(\frac{\Delta}{\Delta + \gamma} \right) \left(\frac{R_s}{\lambda} \right) - 0.12$	[89]
56		McGuinness-Bordne	$ET_o = (0.00597T_{\text{mean}} - 0.0838) R_s$	[90]
57		Stephens	$ET_o = (0.0158T_{\text{mean}} + 0.09) \cdot \frac{R_s}{\lambda}$	[91]
58		Stephens and Stewart	$ET_o = (0.0148T_{\text{mean}} + 0.07) \cdot \frac{R_s}{\lambda}$	[92]
59	Mass Transfer Based	Dalton	$ET_o = (3.648 + 0.722 u_2)(e_s - e_a)$	[40]
60		Meyer	$ET_o = (3.75 + 0.503u_2)(e_s - e_a)$	[41]
61		Rohwer	$ET_o = (3.3 + 0.891 u_2)(e_s - e_a)$	[93]
62		WMO	$ET_o = (1.298 + 0.934 u_2)(e_s - e_a)$	[94]
63		Brockamp & Wenner	$ET_o = 0.5435u_2^{0.456}(e_s - e_a)$	[95]
64		Mahringer	$ET_o = (0.286 (u_2)^{0.5}) (e_s - e_a)$	[96]
65		Penman	$ET_o = 0.376(e_s - e_a) u_2^{0.76}$	[42]
66		Modified Penman	$ET_o = (2.625 + 0.000479u_2)(e_s - e_a)$	[97]
67		Romanenko	$ET_o = 0.00006(100 - RH_{\text{mean}})(T_{\text{mean}} + 25)^2$	[43]
68		Modified Romanenko	$T_o = 4.5 \left(1 + \frac{T_{\text{mean}}}{25} \right)^2 \left(1 - \frac{e_a}{e_s} \right)$	[80]
69		Albrecht	$ET_o = (1.005 + 2.97 u_2)(e_s - e_a)$	[98]
70		Trabert	$ET_o = 0.3075u_2^{0.5}(e_s - e_a)$	[99]
71		Szasz	$ET_o = 0.0053(T_{\text{mean}} + 21)^2(1 - RH_{\text{mean}}/100)^{\frac{2}{3}} \cdot f(u)$ $f(u) = 0.0519u_2 + 0.905$	[100]

Table 2: Variables abbreviation used in Table 1

Notation	Description	Unit
ET_o	reference evapotranspiration	mm/day
Δ	the slope of saturation vapor pressure curve at mean air temperature	kPa/°C
γ	psychrometric constant = 0.054	kPa/°C
λ	latent heat of vaporization = 2.45	MJ/kg
R_n	net radiation	MJ/m ² -day
G	soil heat flux density ≈ 0	MJ/m ² -day
u_2	daily mean wind speed at 2m height	m/sec
e_a	saturation vapor pressure	kPa
e_s	actual vapor pressure	kPa
K_{time}	unit conversion equal to 86400 for ET in mm/day	
ρ_a	mean air density at constant pressure	kg/m ³
c_p	specific heat of the air	MJ/kg °C
r_s	bulk surface resistance	s/m
r_a	aerodynamic resistance	s/m
p	Mean annual percentage of daytime hours = 0.274	
T_{mean}	mean daily temperature	°C
T_{max}	daily maximum temperature	°C
T_{min}	daily minimum temperature	°C
R_s	solar radiation	MJ/m ² -day
R_a	extraterrestrial radiation	MJ/m ² -day
RH_{mean}	Mean daily relative humidity	%
RH_{max}	daily maximum relative humidity	%
RH_{min}	Daily minimum relative humidity	%
z	elevation above mean sea level	m
Φ	Latitude	degrees
ET_{pan}	Pan evapotranspiration	mm/day
n/N	relative sunshine duration	
T_{dew}	dew point temperature	°C

Empirical models have historically played a significant role in the estimation of reference evapotranspiration. Studies by Kashyap and Panda. [101], Landaras et al. [102], Sabziparvar et al. [103], and others investigated several equations for reference evapotranspiration estimation. DehghaniSanij et al. [104] compared various models with lysimeter data and underscored the reliability of the Penman-Monteith model. Additionally, in different climatic conditions, researchers found specific models to be more suitable: the Turc [82] model for humid conditions [83], the Abtew model for various global applications [105], [106], [107], [108], and the Priestly–Taylor, Hargreaves-Samani, and Makkink models for arid regions, semi-arid regions, and humid regions, respectively [67], [109], [110], [111]. This performance of different models in specific environments is due to; Climate and weather patterns (models that account for temperature and humidity such as the Penman-Monteith method tend to perform better in temperate climates with moderate temperature and humidity levels. Models that incorporate radiation and sunshine data such Hargreaves-Samani method tend to perform better in regions with high solar radiation and sunshine hours. Models that account for wind and precipitation patterns such as the FAO-56 method tend to perform

better in regions with high wind speeds and precipitation variability), Vegetation, and land use (Crop types and growth stages: FAO-56 performs better in agricultural regions with diverse crop types and growth stages and Vegetation density and height), Topography and soil (Elevation and slope, soil type and moisture), Data availability and quality, Model complexity and Calibration, and Regional and local factors

Forecasting of Reference Evapotranspiration:

Accurate forecast reference evapotranspiration is pivotal for crop planning, irrigation scheduling, and water allocation. Several methods have been employed to forecast reference evapotranspiration which are broadly classified;

Statistical Models: Regression models have been used to forecast reference evapotranspiration based on historical meteorological data [112].

Artificial Neural Networks (ANNs): ANNs have shown promise in forecasting reference evapotranspiration using historical data by learning complex relationships between reference evapotranspiration and meteorological parameters [113].

Physical Models: Physical models such as the Penman-Monteith equation can be employed for forecasting by utilizing the current or forecasted meteorological data. These models require detailed meteorological data and may be computationally intensive [114].

Machine Learning: Machine learning algorithms, trained on historical meteorological data forecast reference evapotranspiration. While combining statistical and physical models, these algorithms provide accurate forecasts [115].

Autoregressive Integrated Moving Average (ARIMA) models have historically been used for time series forecasting [116], [117], [118]. However, recent years have seen increased interest in ANN models [119], [120], [121], [122], [123], [124], [125]. Moreover, advanced techniques in deep learning hold promise for the future of reference evapotranspiration forecasting [125], [126], [127], [128]. The summary of the published articles that used different models for a forecast of reference evapotranspiration is given in Table 3.

The studies presented in Table 3 reflect the capability of ANN and machine learning models in the forecast of reference evapotranspiration. The ANN models show greater performance compared to other conventional models. Deep learning models, particularly RNN and LSTM are well suited for handling temporal dependencies by distributing representations, temporal hierarchies, automatic feature learning, and flexibility and adaptability. Traditional statistical models like ARIMA and VAR rely on the explicit modeling of temporal relationships using predefined equations. These models have limitations of assuming stationarity, limited flexibility, and require manual feature engineering. Deep learning models yield outstanding performance in the forecast of reference evapotranspiration. However, Some of the ANN and most recent deep learning techniques such as Recurrent Neural Networks (RNN), Convolutional Neural Networks (CNN), Generative Adversarial Neural Networks (GANN), Gated Recurrent Units (GRU), Transformers, Temporal-Convolutional Neural networks (TCNN), Long-Short Term Memory (LSTM), and hybrid models have very limited utilization in forecast reference evapotranspiration.

Artificial Neural Network (ANN), Multivariate Relevance Vector Machine (MVRVM), Support Vector Machine (SVM), Gaussian Process Regression (GPR), Long short Term Memory (LSTM), Seasonal Auto-Regressive Integrated Moving Average (SARIMA), Group Method of Data Handling (GMDH), One-Dimensional Convolutional Neural Network (1D-CNN), Random Forest (RF), Multiple Linear Regression (MLR), Adaptive Neuro Fuzzy Inference System (ANFIS), Numerical Weather Prediction (NWP), Auto-Regressive Integrated Moving Average (ARIMA), Multilayer Perceptron (MP), Generalized Feed Forward (GFF), Linear Regression (LR), Probabilistic Neural Network (PNN), Genetic Programming (GP), Temporal Convolution Neural Network (TCNN), Adaptive Boosting Machine Learning (ABML), Vector Auto-Regression (VAR), K-Nearest Neighbor (K-NN),

Deep Neural Network (DNN), Support Vector Machine (SVM), Least Square SVM (LSSVM), Genetic Expression Programming (GEP), Moving Average (MA), Auto-Regression (AR), Auto-Regressive Moving Average (ARMA), Generalized Regression Neural Network (GRNN), Auto Encoder-Decoder Bidirectional LSTM (AED-BiLSTM)

Table 3: Forecast of Reference Evapotranspiration

S. No	Forecast	Model(s)	References
1	Monthly	ANN	[129]
2	Daily	ANN and MVRVM	[130]
3	Daily	GPR and Wavelet-GPR	[131]
4	Monthly	SARIMA, GMDH, and SVM	[132]
5	Daily	LSTM, 1D-CNN, CNN-LSTM, ANN, and RF	[133]
6	Monthly	SVM, ANFIS, ANN, MLR	[134]
7	One-day Ahead	ANN	[119]
8	Daily	NWP	[135]
9	Summer ET _o	Multiple Global NWP	[136]
10	Daily	NWP	[137]
11	Daily	NWP	[138]
12	Daily ET _o	NWP	[139]
13	Weekly	ARIMA and ANN	[117]
14	Weekly	ANN	[123]
15	Near Future	GFF, LR, MP, and PNN	[124]
16	Monthly	GP, SVM, ANN, and SVM-Wavelet	[121]
17	Daily	VAR	[140]
18	Daily	SVM, ANN	[141]
19	Daily	k-NN, ANN, and ABML	[142]
20	Daily	DNN, LSTM, TCNN	[143]
21	Daily	SVR	[144]
22	Monthly	SVM and GEP	[145]
23	Daily	GEP and ANN	[146]
24	Daily	RF and GEP	[147]
25	Daily	Univariate AR and MA	[148]
26	Daily	LSTM	[149]
27	Daily	AR, MA, ARMA, ARIMA, LSSVM, ANFIS, and GRNN	[150]
28	Weekly	AED-BiLSTM	[151]

The accuracy of forecast reference evapotranspiration is evaluated by using different models and techniques. Ballesteros et al. [152], used the Penman-Monteith equation as a reference methodology and compared the forecast reference evapotranspiration from 2011 to 2012 by using ANN and Hargreaves-Samani equation. The Hargreaves-Samani model is also applied by Luo et al. [139], for daily forecast reference evapotranspiration with an allowable degree of accuracy. Pelosi et al. [138] used Hargreaves-Samani and Priestly-Taylor methods

for daily forecast reference evapotranspiration with a short-term lead time of 1 to 5 days using Numerical Weather Prediction (NWP) outputs in Southern Italy. Yang et al. [153], utilized the public weather forecast for size locations of wide climatic ranges in China to forecast 7 days daily reference evapotranspiration.

Short, Medium, and Long-Term Forecast Reference Evapotranspiration:

Some studies forecasted the reference evapotranspiration through meteorological data for different periods [48], [154], [155]. [137], used two basic and primary approaches, the time series and ANNN, to forecast the medium to long-term reference evapotranspiration by using meteorological data. Tracy et al. [156], utilized metrological data and forecasted the yearly reference evapotranspiration by using the ARIMA model. Marino et al., [156], recommended the seasonal ARIMA model for accurate monthly forecast reference evapotranspiration. While assessing short-term to medium irrigation decisions, some investigators revealed that forecasting daily and weekly reference evapotranspiration is more effective in decision-making [102], [129], [157], [16]. Therefore, the trend is turned from a long-medium-term range to a medium-short-term range, because the forecast of daily reference evapotranspiration is more helpful in scheduling short-term irrigation activities. A modified daily mean model is presented by [152], [123], and [158] to forecast the daily reference evapotranspiration. Similarly, [137], forecasted the daily reference evapotranspiration with good results for 9 days lead time by using the Penman-Monteith equation.

Forecast of Reference Evapotranspiration by Satellite Observations:

In past studies, different data sources were utilized to forecast reference evapotranspiration. Most of the researchers have utilized the ground-based observed meteorological data to quantify the reference evapotranspiration and project it for the future [111], [135], [141], [143], [159], [160]. However, over the large areas, the meteorological variables are nonhomogeneous which have different effects on reference evapotranspiration. So, it is important to evaluate the forecast reference evapotranspiration by different models, from different data sources, and for different localities. Several remote sensing models have been developed for estimating reference evapotranspiration, including:

SEBAL (Surface Energy Balance Algorithm for Land) uses the combination of satellite data (e.g., Landsat, MODIS) and meteorological data to estimate the surface energy balance components, including reference evapotranspiration [161]. METRIC (Mapping Evapotranspiration at High Resolution with Internal Calibration) is another popular remote sensing model for estimating reference evapotranspiration [162]. It uses a combination of satellite data (e.g., Landsat, ASTER) and meteorological data to estimate reference evapotranspiration at high spatial resolution. MOD16 (Moderate Resolution Imaging Spectroradiometer Global Evapotranspiration) is a global reference evapotranspiration product derived from MODIS satellite data [163]. It provides daily reference evapotranspiration estimates at 1 km spatial resolution, which can be useful for regional and global water resources management.

Optical remote sensing uses visible and infrared radiation to estimate reference evapotranspiration with high spatial resolution (up to 1-2 meters). Some common optical remote sensing methods for estimation of reference evapotranspiration include Vegetation Indices, surface temperature, and albedo. While microwave remote sensing uses longer wavelengths (around 1-10cm) to estimate reference evapotranspiration. Soil moisture and vegetation water content estimates from microwave radiation are related to evapotranspiration.

Satellite-based estimation of reference evapotranspiration has several advantages e.g., high spatial resolution and coverage, ability to estimate reference evapotranspiration, and potential for real-time monitoring. While compared to ground-based meteorological data, all satellite-based data parameters have some level of uncertainty that needs to be considered.

Limitations are experienced in obtaining accurate reference evapotranspiration from remote sensing techniques due to the temporal resolution of satellite overpasses and/or gaps in images due to cloud covers [164]. Therefore, the comparison of reference evapotranspiration from satellite data with actual reference evapotranspiration needs proper attention. Satellite observations are combined with forecast reference evapotranspiration to schedule and control the irrigation by [165], [166], and [166]. The best results were observed by [167], while comparing the forecast reference Evapotranspiration from FRET and observed reference evapotranspiration.

Reliability of Forecasted Reference Evapotranspiration:

The reliability of estimated reference evapotranspiration from forecasted meteorological data must be analyzed with actual meteorological data for effective decisions in water management. The author in [168], compared the results to evaluate the estimated reference evapotranspiration, irrigation scheduling, and yield from forecasted meteorological data with observed data. Another study [160], separately considered the accuracy by using the past and forecasted data inputs, to evaluate the reliability of short-term weather forecasts. In contrast to the studies presented in Table 3, some studies have utilized the forecasted meteorological data to forecast reference evapotranspiration, such as public weather forecasts [124], [135], [137], [153]. However, it is not clear from the previous studies which scenario is more accurate to forecast reference evapotranspiration, either using future forecasted meteorological data or using the historical reference evapotranspiration data for future projections.

Sensitivity in Forecasting Reference Evapotranspiration:

According to [152], for high crop water demanding periods the accurate forecast reference evapotranspiration is of utmost importance. Because the poor quantification of crop water requirement can affect the crop yield and/or water utilization. Scheduling the next irrigation activity based on the previous week's reference evapotranspiration without taking into account the variation in weather conditions from one week to the next, can result in over and/or under irrigation [169]. In this context, accurate forecast reference evapotranspiration is important for improving irrigation efficiency. Usually, the output of the forecasting models is available to scientists and researchers but not to the end users, farmers, and irrigation managers. So, more precise techniques and models are needed to accurately forecast reference evapotranspiration to enhance the efficient use of water.

Conclusion:

Evapotranspiration is a critical parameter for various applications, including irrigation planning, and climatological and hydrological studies. However, the high cost associated with installing and maintaining physical networks for measuring evapotranspiration, limited data is available in many regions worldwide. Additionally, various local, regional, and global factors affect reference evapotranspiration, adding uncertainty to its estimation. This uncertainty is further compounded by the influence of climatic factors.

Over the past century, empirical and physically based models have been used to estimate reference evapotranspiration, with the Penman-Monteith FAO56 model widely accepted due to its superior performance. However, the limited availability of complete climatic data needed for this model has led to the use of temperature-based models which use fundamental meteorological data collected at most weather stations.

To meet the need for accurate crop water requirement estimation, forecast reference evapotranspiration is indispensable. Reference evapotranspiration is a complex, non-linear, and dynamic phenomenon. Researchers have employed diverse techniques, from traditional statistical models to more recent deep learning models, to forecast reference evapotranspiration. Deep learning models show remarkable promise in forecast reference evapotranspiration. Nonetheless, the hydrology and climatology fields are yet to fully explore

the potential of the most recent deep-learning models, and future research should delve into these untapped possibilities.

Challenges and Future Directions:

The methods and models presented in this study have significantly enhanced our understanding of reference evapotranspiration, but challenges remain. Incorporating real-time data, addressing model accuracy, and considering the implications of climate change are avenues ripe for exploration. Further research in these areas promises to refine existing models and develop novel, robust forecasting techniques.

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